condition, and their original character so completely changed, that it becomes a matter of exceeding difficulty to place them in their proper geological sequence. Lithologically, they bear no resemblance to the typical Laurentian, nor yet do they possess much in common with the Huronian or intermediate system of Mr. Murray. Moreover, as will presently appear, they are overlaid unconformably by strata, having many characteristics of the latter.

Resting upon these in the valley stretching inland from the head of Clode Sound, a much less altered set of sedimentary strata are found. These consist of red and gray silicious sandstones and grits or fine conglomerates, with several arenaceous slaty divisions, bearing a most striking lithological resemblance to the Signal Hill sandstones near St. John's. The greenish-gray sandstones at the base are not so well represented as the redder strata, division f and g. of Mr Murray's section. Some portions towards the top are more slaty and somewhat micaceous, and would seem to mark an upward extension of the same series not brought in on Signal Hill. The whole stretch across the head of the Sound from the valley of the S. W. River to Salmon River on the northern side, and form a set of wave-like undulations. One or two small greenstone dykes are seen to cut these strata at right angles, but apparently do not cause much extra disturbance. On the northern side of this trough the red sandstones suddenly become more disturbed, are highly tilted and at length distinctly overturned; while at the same time their character has greatly changed, and they become altered to a pale, pinkish quartzite, possessing a remarkably pretty, roseate hue.

A deep depression within a couple of hundred yards of Salmon River Ridge marks the line of contact of this set of rocks with an entirely distinct and much more metamorphosed series. Evidently a fault occurs at this junction, which appears to follow the course of the depression just alluded to. In the bottom of this depression low out-crops of a fine-grained, compact quartzite occur, dipping N. 20, W. angle 66 degrees. These are succeeded by chlorite schists of a dark-green color merging into grayish-purplish and pale-yellowish nacreous or talcose slates. The pale-colored bands in particular are highly talcose, exhibiting scales of pure, yellowish-green talc between the layers of stratification, as well as on the surfaces of the cleavage plains. Some of these slates are very fissile, having