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glacier may still have continued to hold some importance in the foot-hill region, but the abundant supply of well rounded gravels, with other circumstances, renders it probable that the Rocky Mountain glaciers generally had become strictly local and relatively insignificant.

If it may thus be assumed that the higher terraces and traveled gravels of the Porcupines are approximately contemporaneous with the upper boulder-clay, all the lower and later terraces and gravel plains may be regarded as marking stages in the subsidence of this water-level from its maximum height of 5,300 feet. These, it has already been noted, are usually not strongly impressed, and there is no evidence that the subsidence was arrested long, except at one stage, which is that spoken of in the report of 1882-'84 as being at about 4,200 feet. Further examination appears to show that the terraces referable to this particular stage slope up gradually in the foothills and on approaching the mountains to a maximum height of about 4,500 feet, from which it may be argued that from the last mentioned height the water lowered its level gradually to one of about 4,200 feet, while new material was constantly being washed down by rivers from the mountains. A later and still lower, though less important, period of arrest seems to be marked by the gravel plain near Macleod at about 3,200 feet.

The first mentioned line of relative stability appears to be equally well marked in the southern portion of the region, about Waterton lake and the Oldman river, and in the northern, in the Bow valley, leading to the suggestion that the irregular uplift of the earlier stages of recovery had been succeeded along the base of the mountains by one in which further change of level occurred throughout uniformly, as compared with the actual heights of the surface found in the same region today, or with isobases changed in direction and parallel to the trend of the mountains. This later uplift may have continued, with the stranding of large boulders near the water-line from time to time, until this part of the plains reached its present condition and slope.

There is, however, some good evidence to show that in postglacial times a renewed or continued southern uplift took place. This is derived from the changes in the course of streams and slopes of their valleys, but cannot be entered into in this paper.\*

In this connection I may digress so far as to mention that there is a somewhat notable correspondence between the higher levels of terraces on both sides of the Rocky mountains and continental watershed. It is found in the southern part of the interior plateau of British Columbia

\* Report of Progress, Geol. Survey of Canada, 1882-'84, p. 150 C; Annual Report, Geol. Survey of Canada, vol. I (n. s.), p. 75 C; Physiographical Geology of the Rocky Mountain region in Canada, Trans. Royal Soc. Canada, vol. viii, sec. 4, p. 63.