

those of the Queen-Charlotte Islands. Similar erratics appear to characterize in greater or less abundance the whole of the low country above described, but are not found about the heads of the south-western extremities of Masset Inlet.

It has previously been shown that at the time when the Strait-of-Georgia glacier began to diminish the sea must have stood considerably higher in relation to the land than at present, and the glaciated rock surfaces about Victoria and Nanaimo no sooner appeared from beneath the glaciers than they were covered by deposits holding marine shells. Such must have been the state of affairs also in the Queen-Charlotte Islands; and to this time are doubtless to be referred the clay and sand deposits of the low north-eastern part of Graham Island above described. The material of these must have been supplied from the glaciers of the islands themselves, but added to also (as the nature of the boulders proves) by the *débris borne* on floating ice from the larger glaciers of the mainland, the sea levelling and spreading abroad the material, and preventing the formation of any well-marked terminal moraines by the island glaciers. The rocky beds of the fiords and Masset-Inlet expansions must have been shaped to some extent by the ice; but the absence of drift material from their areas, and especially of the erratics derived from the mainland, are, with their situation, good reasons for supposing that they mark the regions last covered by glacier-ice, and from which it eventually retreated with some rapidity, leaving the hollows formerly occupied by it to become first inlets, and then, with increasing elevation, in some instances lakes.

The exceptional case which seems to show the impingement on the Queen-Charlotte Islands of ice not produced on them was found on the north coast on the little islands lying outside the entrance to Masset Inlet; but it is probable that similar traces might be found by search in additional localities in this vicinity. Wider exposures of basalt a few feet above high-water mark here show very heavy though somewhat worn glaciation in a direction S. 10° E., or N. 10° W., but probably the former. The depth and parallelism of the grooving would appear to show that it is glacier work. The mountainous axis of the islands in this their northern part does not exceed in height about 1300 feet, and where nearest is about 15 miles from the locality, while the direction of the marking is not that which would be followed by ice descending from the mountains under any circumstances, being more nearly parallel to than radiant from them. It is, however, just that which ice-masses floating up or down the strait separating the islands from the mainland must have taken, or glacier-ice pushing southward from the long fiords of the Prince of Wales group in Southern Alaska, sixty miles distant. It may, I believe, be attributed with greatest probability to the last-named agent; and in view of the great extension which the glaciers of other parts of the coast must at one time have had, that required for the Prince of Wales group and adjacent channels does not appear excessive.