Professor J. B. Woodworth, is a very useful one, for O, deduced from various stations should be the same for any particular earthquake. This table is due to Dr. Mohorovičić of Agram.

Table 2. This table gives the time interval between the arrival of the P waves and of the PR_1 waves, the latter having been once reflected, hence the time of arrival of the latter will be twice the time a P wave requires to travel the half distance or $\frac{\Delta}{2}$. This table, as well as the next one, table 3, has been directly computed from the preceding one.

Table 3. This table gives the time interval between the arrival of the P waves and of the PR_2 waves, the latter having been twice reflected.

Table 4. This table gives the time interval between the arrival of the P waves and of the PS waves, the path of the latter being one part P waves, the other part S waves, and which may be written too as PR_1S or SR_1P waves.

Table 5. This is the well-known table of Zeissig, based on Wiechert and Zöppritz's values, so useful in determining the distance, Δ, of an earthquake, its epicentre, from the seismological station. It expresses the time-interval between the arrival of the P and S or transverse waves.

Table 6. This table has been compiled from tables 1 and 5, being the sum of the two, but expressed in a smoothed series of whole seconds. For distant earthquakes the horizontal component is weak, often wanting; hence O may be obtained from $S-O$, written too I_S, provided we obtain the distance, Δ, from $L-S$. Knowing O, even approximately, we can now look for P, and frequently identify it then. Applying now $P-O=I_P$, the value of O can be corrected.

Table 7. This table gives the time interval between the arrival of the S waves and of the SR_1 waves, the latter having been once reflected. The derivation is from table 6, and computed similarly to table 2.